












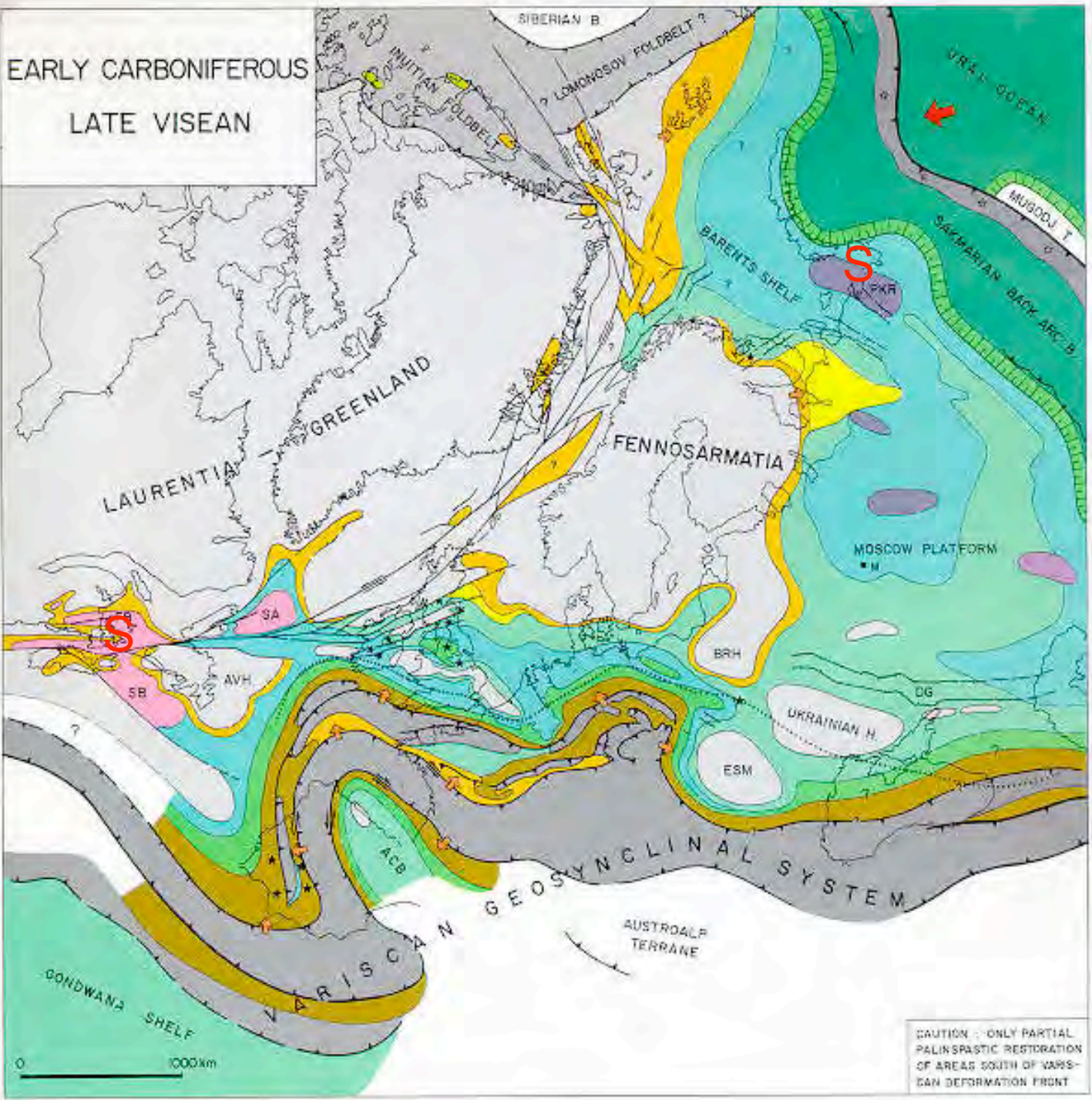


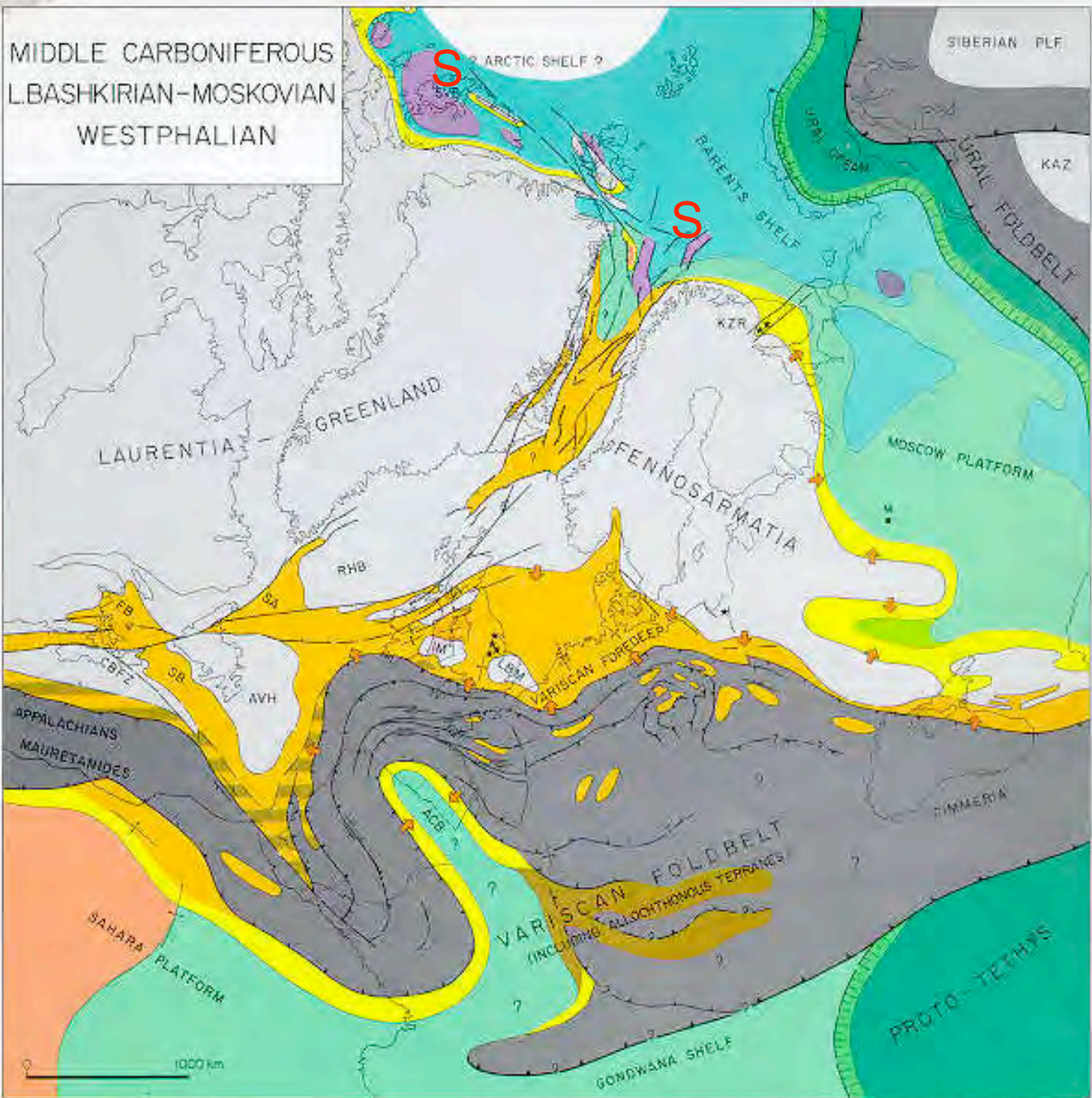
-  MAINLY CONTINENTAL CLASTICS
-  EVAPORITES, CLASTICS AND CARBONATES
-  DELTAIC-SHALLOW MARINE, MAINLY SANDS
-  EVAPORITES AND CARBONATES
-  SHALLOW MARINE, MAINLY SHALES
-  DEEPER MARINE CLASTICS AND/OR CARBONATES
-  SHALLOW MARINE, CARBONATES AND CLASTICS
-  DEEPER MARINE, MAINLY SANDS (FLYSCH)
-  SHALLOW MARINE, MAINLY CARBONATES
-  BASINS FLOORED BY OCEANIC CRUST
-  EVAPORITES AND CLASTICS
-  BASINS FLOORED BY OCEANIC CRUST CONTAINING THICK SEDIMENTS
-  MAINLY EVAPORITES

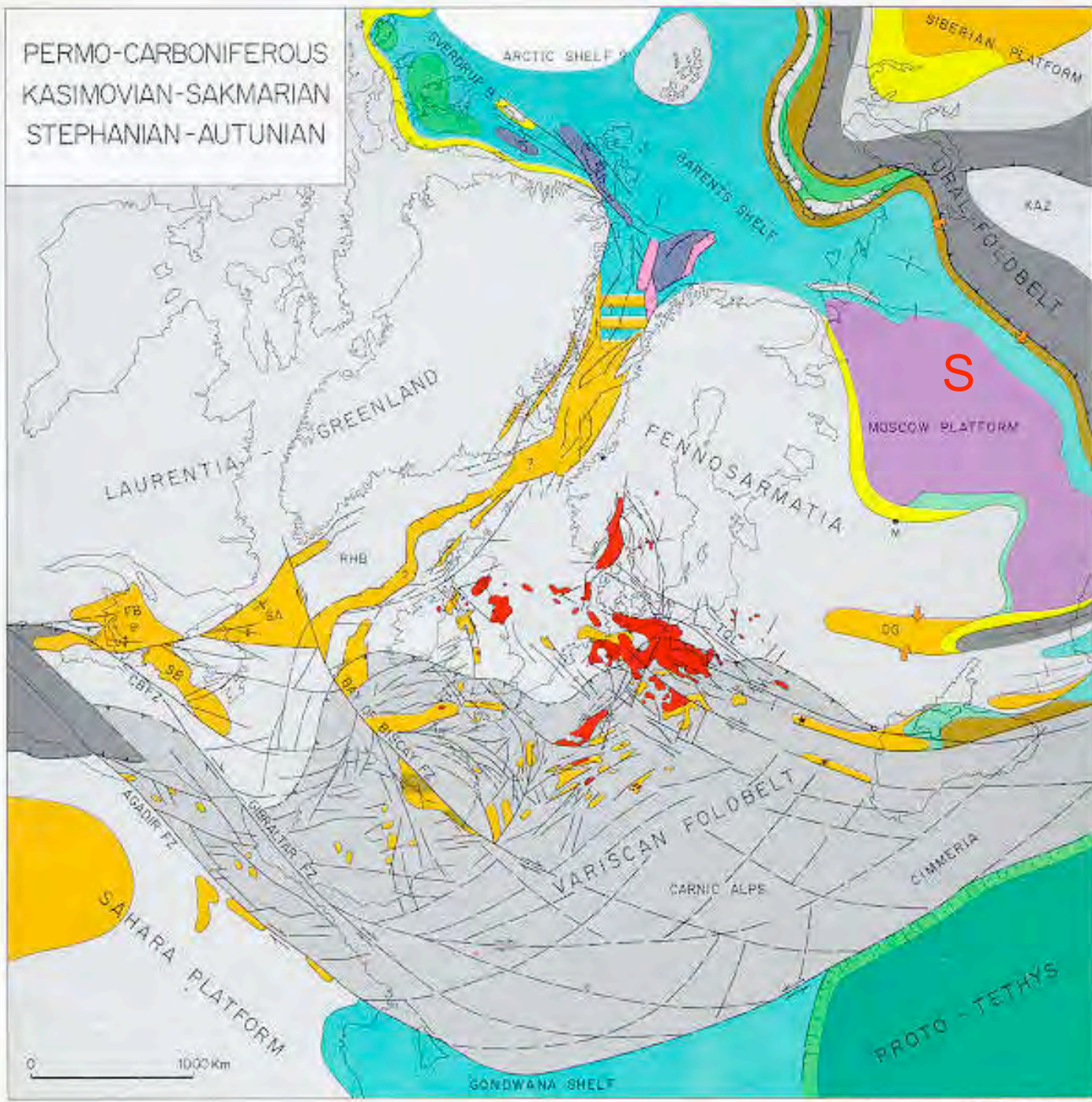
EVOLUTION OF THE ARCTIC-NORTH ATLANTIC AND THE WESTERN TETHYS--A VISUAL PRESENTATION OF A SERIES OF

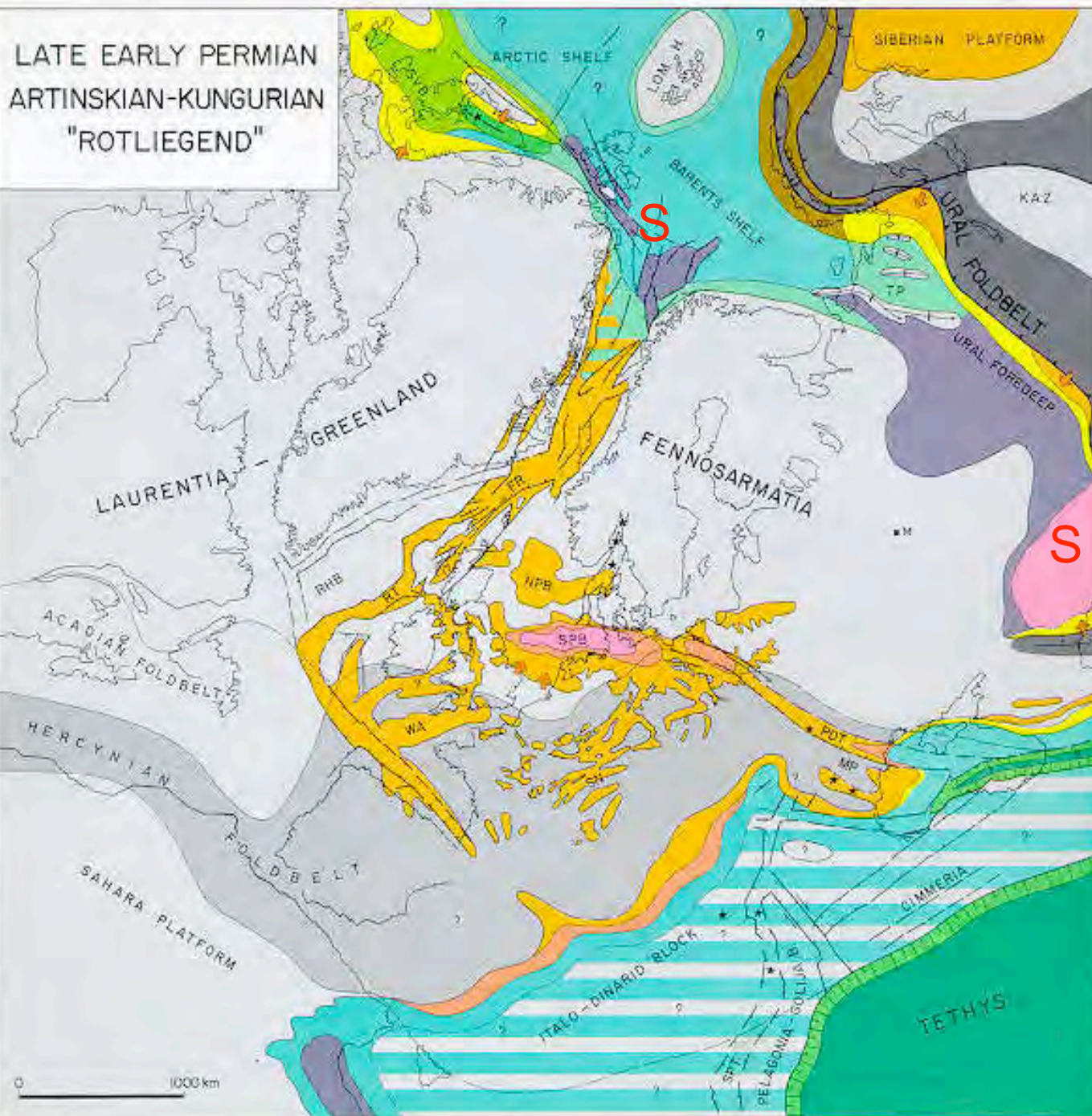
PALEOGEOGRAPHIC-PALEOTECTONIC MAPS*, by Peter A. Ziegler

<http://www.searchanddiscovery.com/documents/97020/>

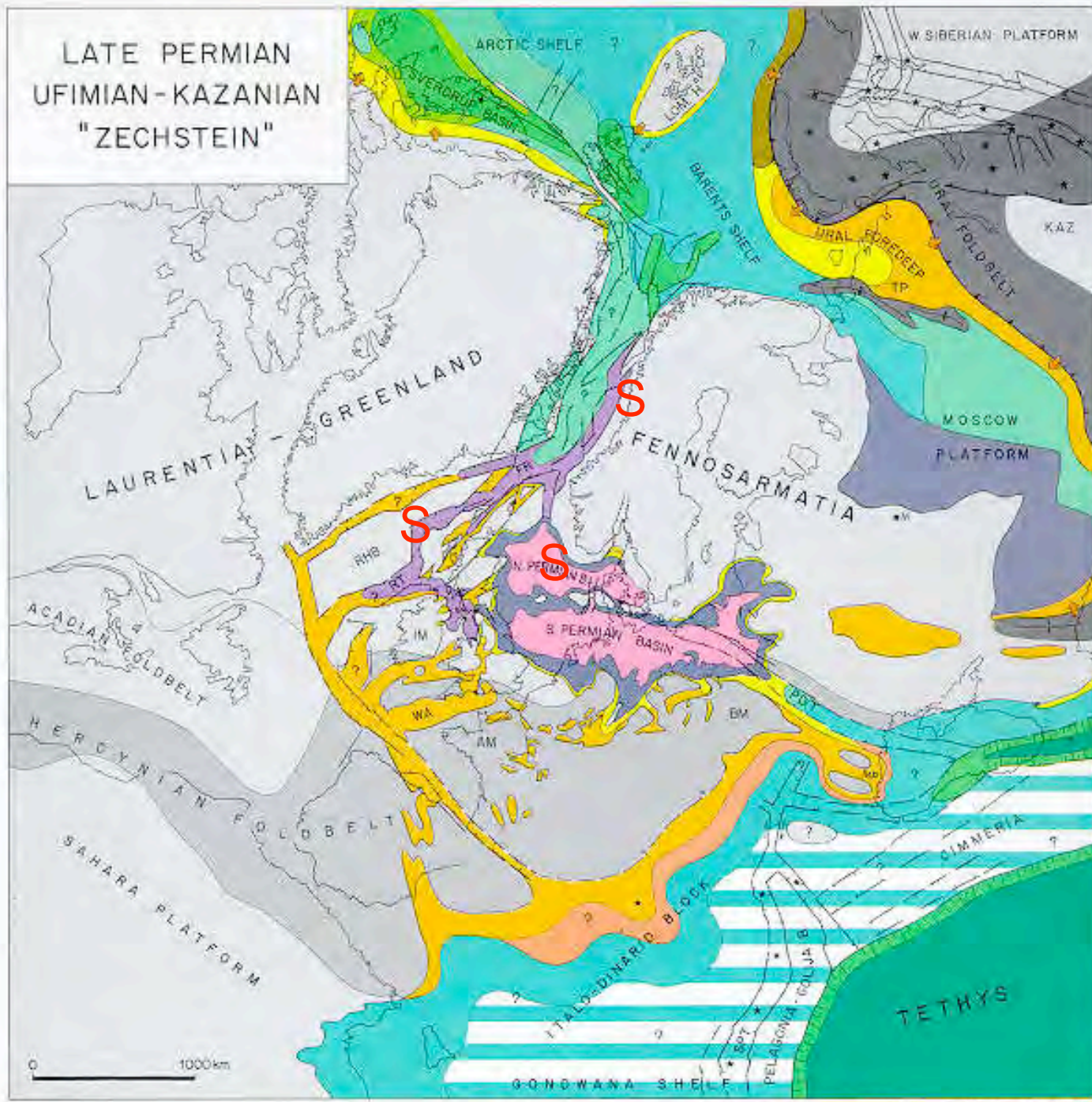




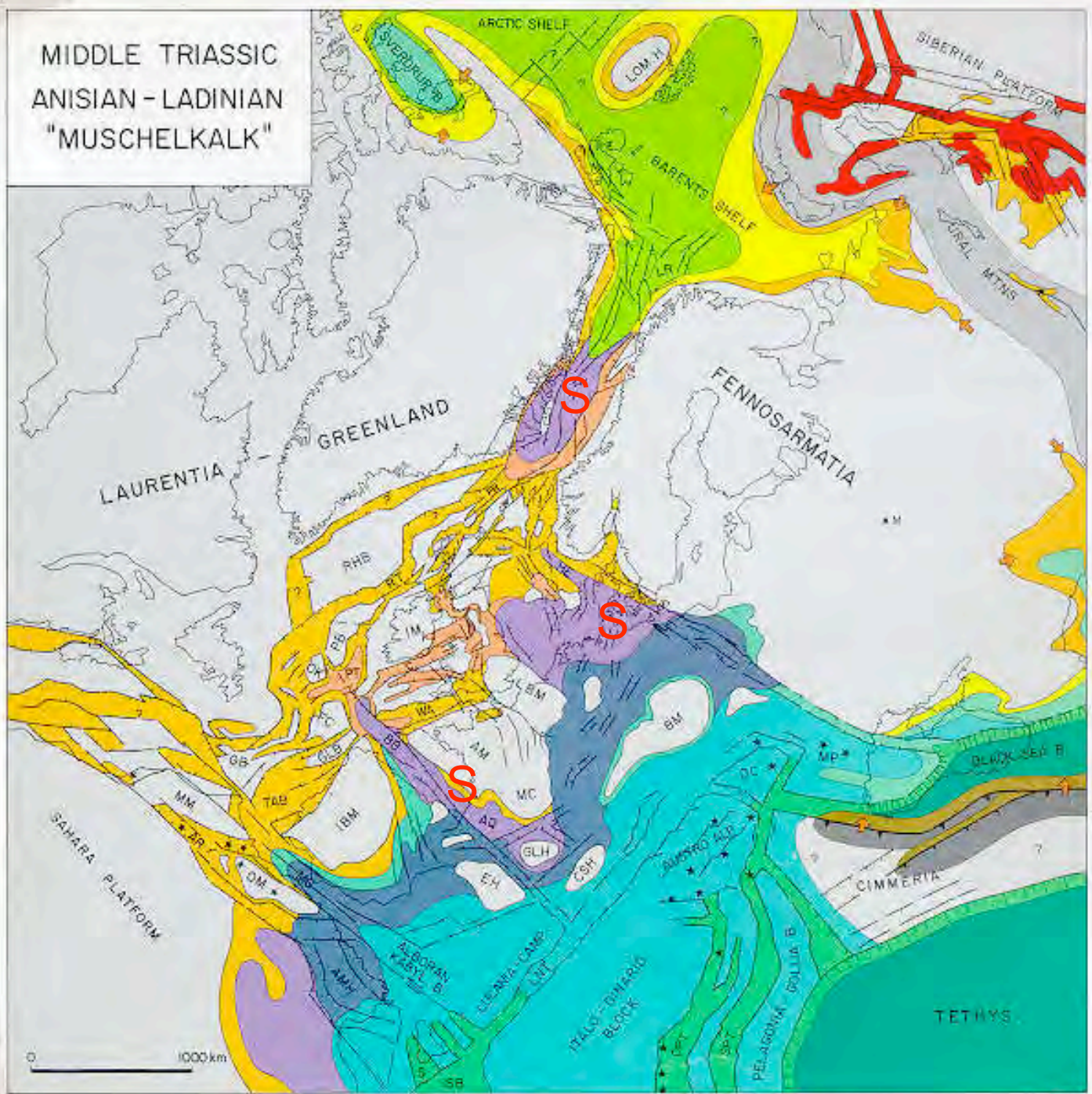


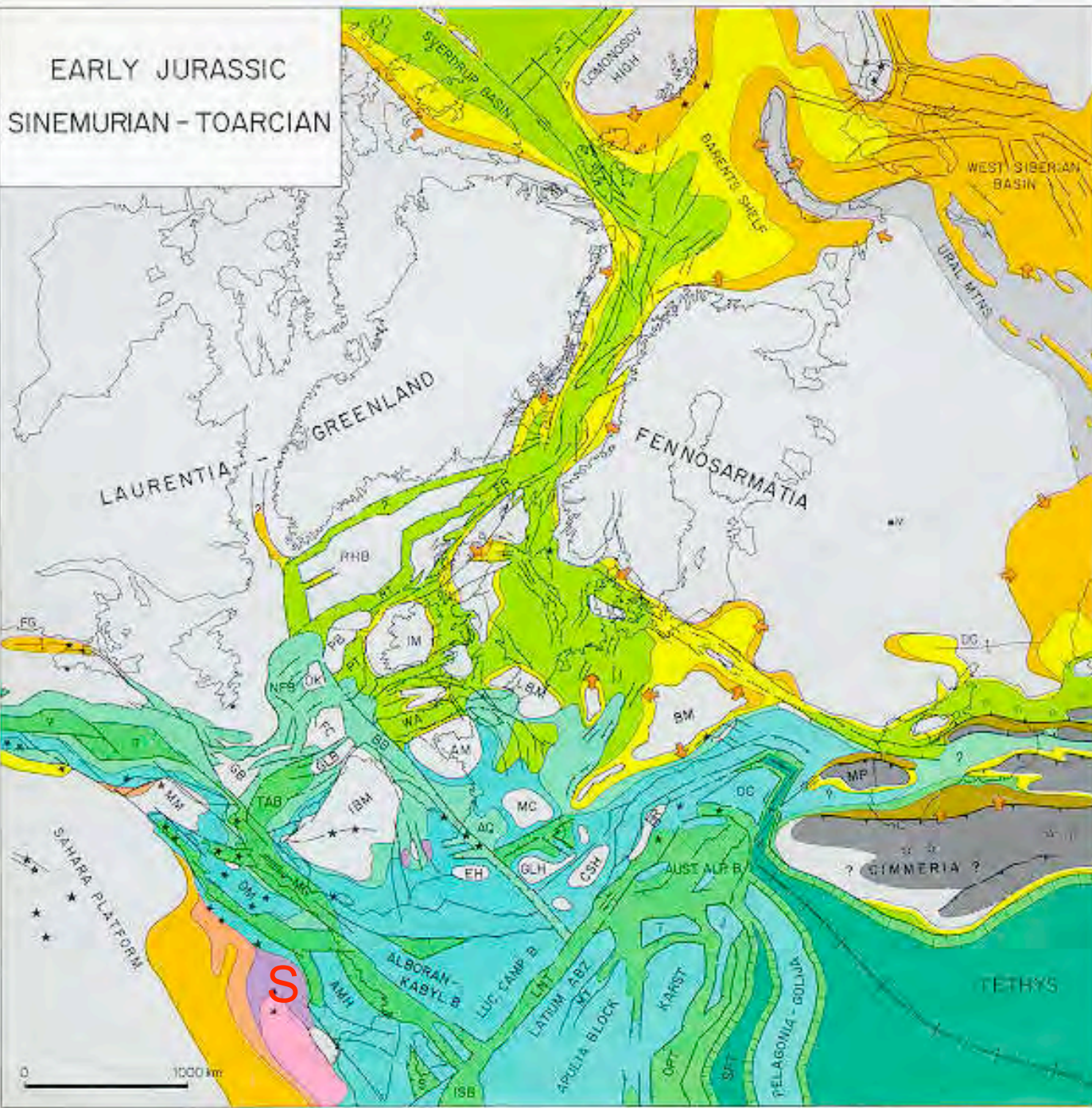


LATE PERMIAN
UFIMIAN-KAZANIAN
"ZECHSTEIN"

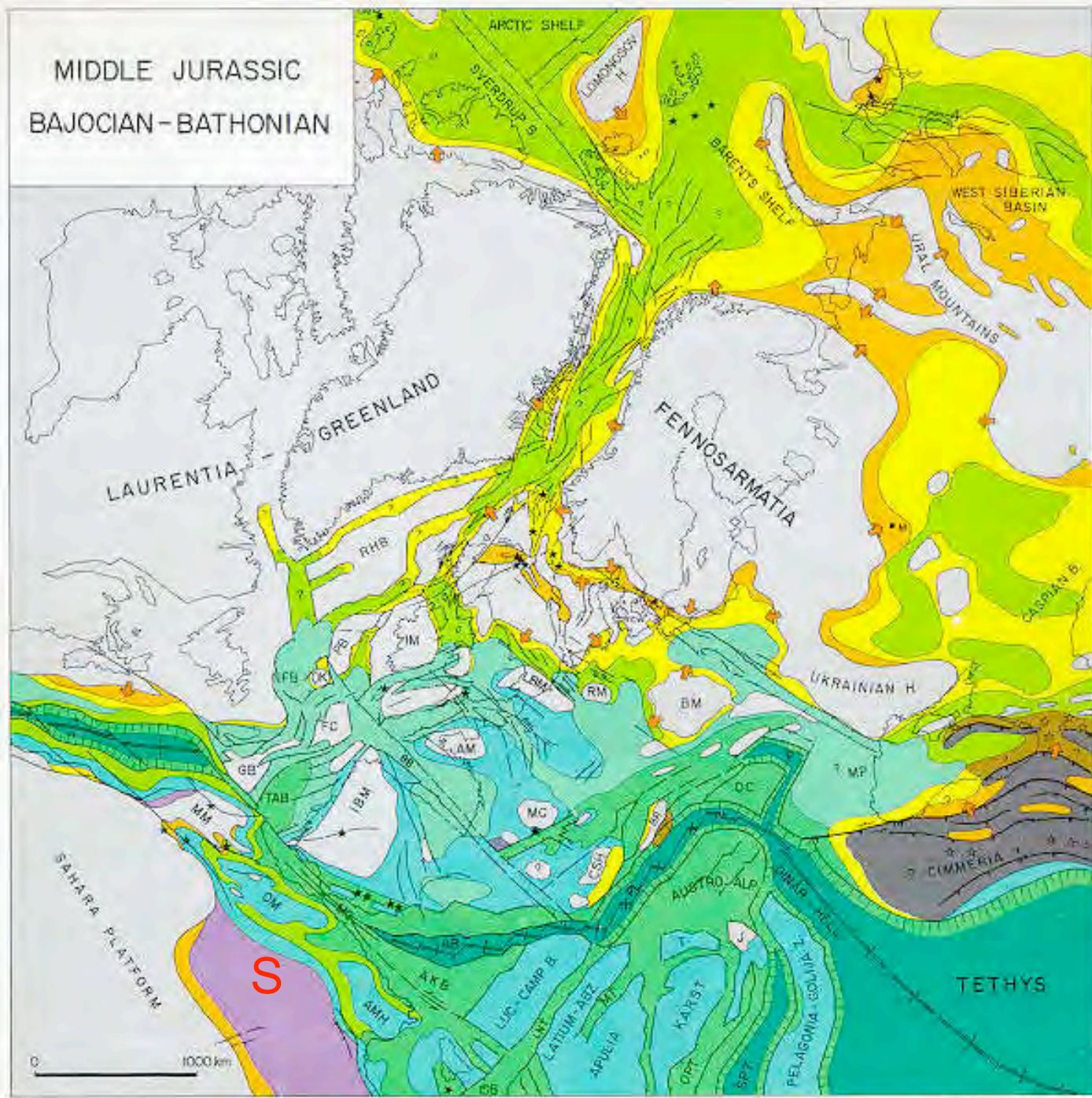


MIDDLE TRIASSIC
ANISIAN - LADINIAN
"MUSCHELKALK"





MIDDLE JURASSIC
BAJOCIAN-BATHONIAN





Great Salt Lake: internally
Drained basin



Trucial coast (UAE) sabhka salt
Pans on margin of Arabian Sea

Two modern evaporite
environments

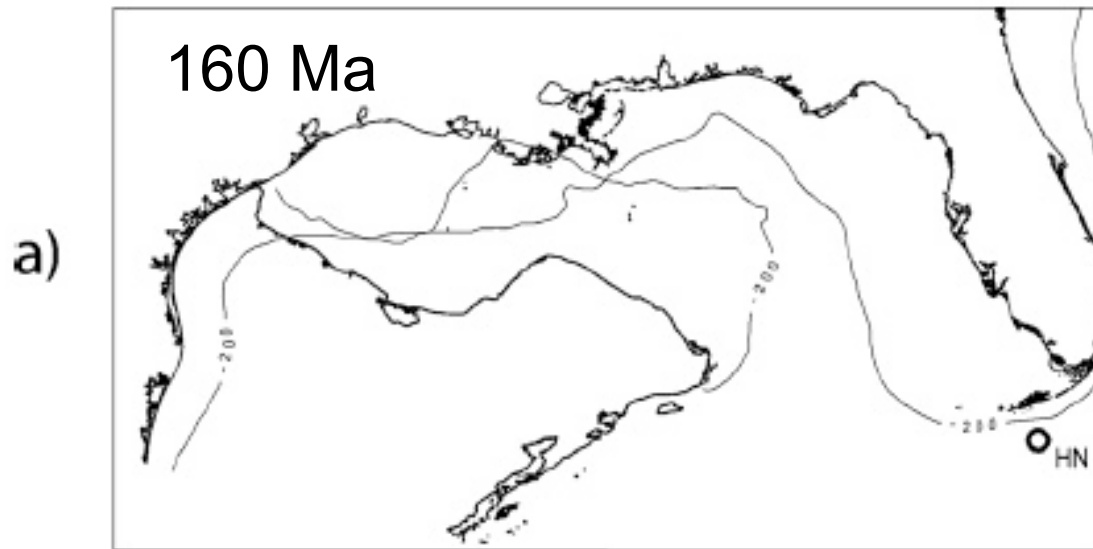
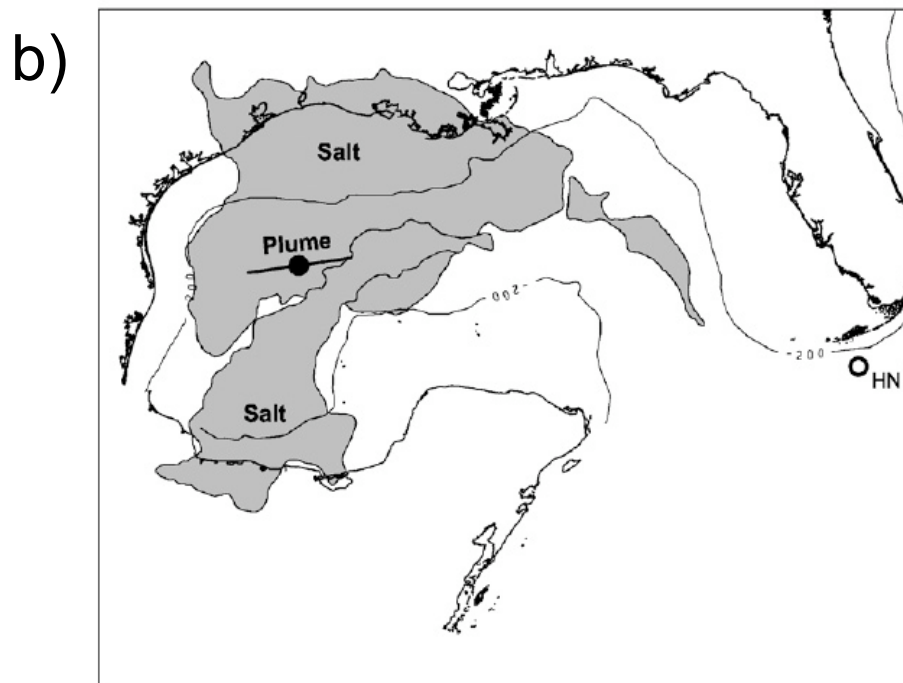
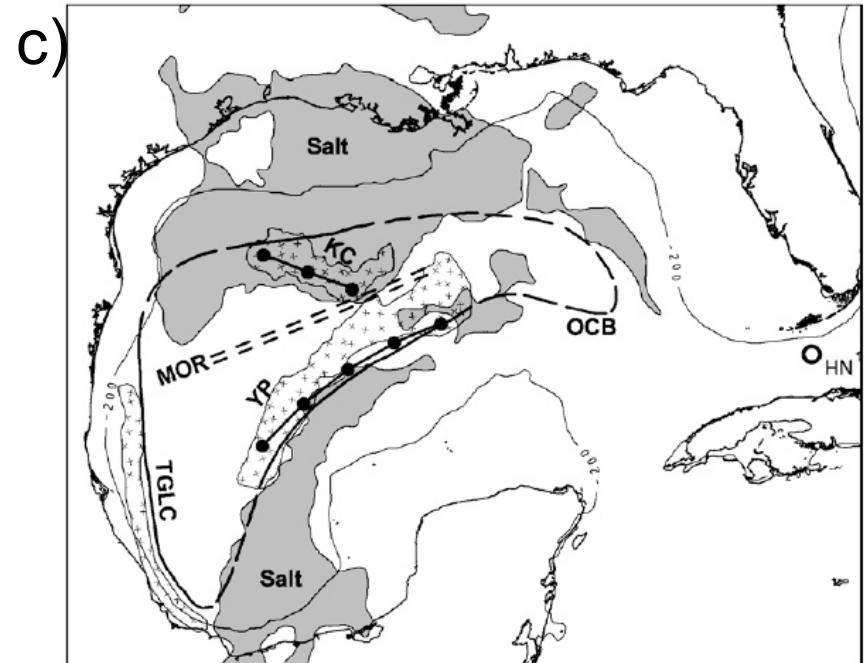


Figure 7. Reconstruction of Gulf of Mexico, 20-m.y. evolution of Yucatan motion. Pole used by Hall and Najmuddin (1994) = HN. (a) Initial position: about 160 Ma (exact age unknown). Yucatan occupies what is the Gulf of Mexico Basin now. Because the Yucatan was probably longer at that time, no gap was present between the peninsula and western Florida (Burke, 1988). (b) 10–12 m.y. coinciding with 22° of rotation and continental crust extension (about 150–152 Ma). Sea-floor spreading began at the end of this time when the plume became active. (c) 20 m.y. and 42° total rotation (adding 20° by rotation of sea-floor spreading), present position achieved (about 140 Ma).



152-150 Ma



~140 Ma

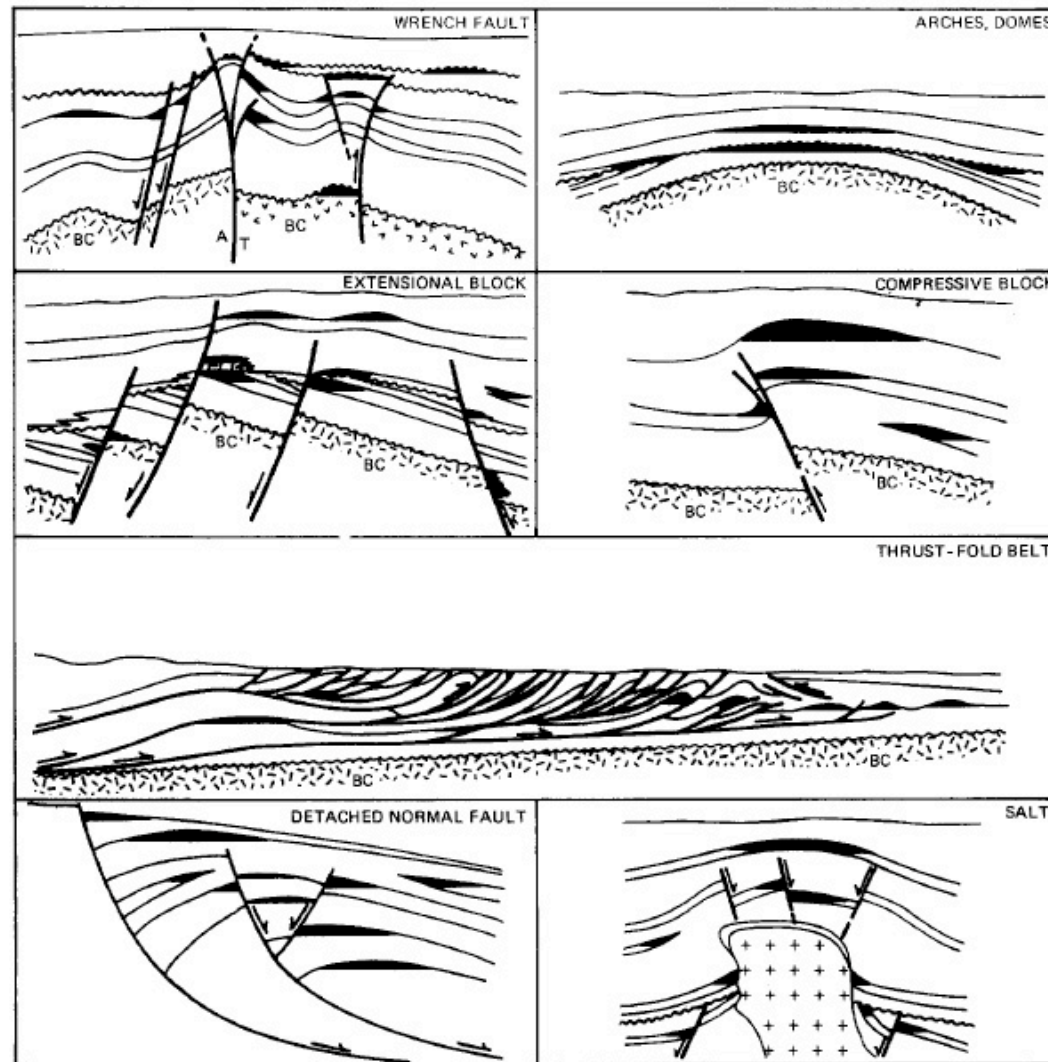


FIG. 1—Schematic diagrams of hydrocarbon traps (black areas) most commonly associated with structural styles of sedimentary basins. Purely stratigraphic type traps and traps associated with basement thrusts are omitted. Salt-related closures modified after Halbouty (1967, Fig. 6). BC, basement complex; T, displacement toward viewer; A, away from viewer.

Grove fault of the southern Illinois basin appears to be a characteristic example (Wilcox et al, 1973). Not all midplate wrench faults developed

style and include most elements that are fundamental to other styles. Wrench assemblages have both compressional and extensional features

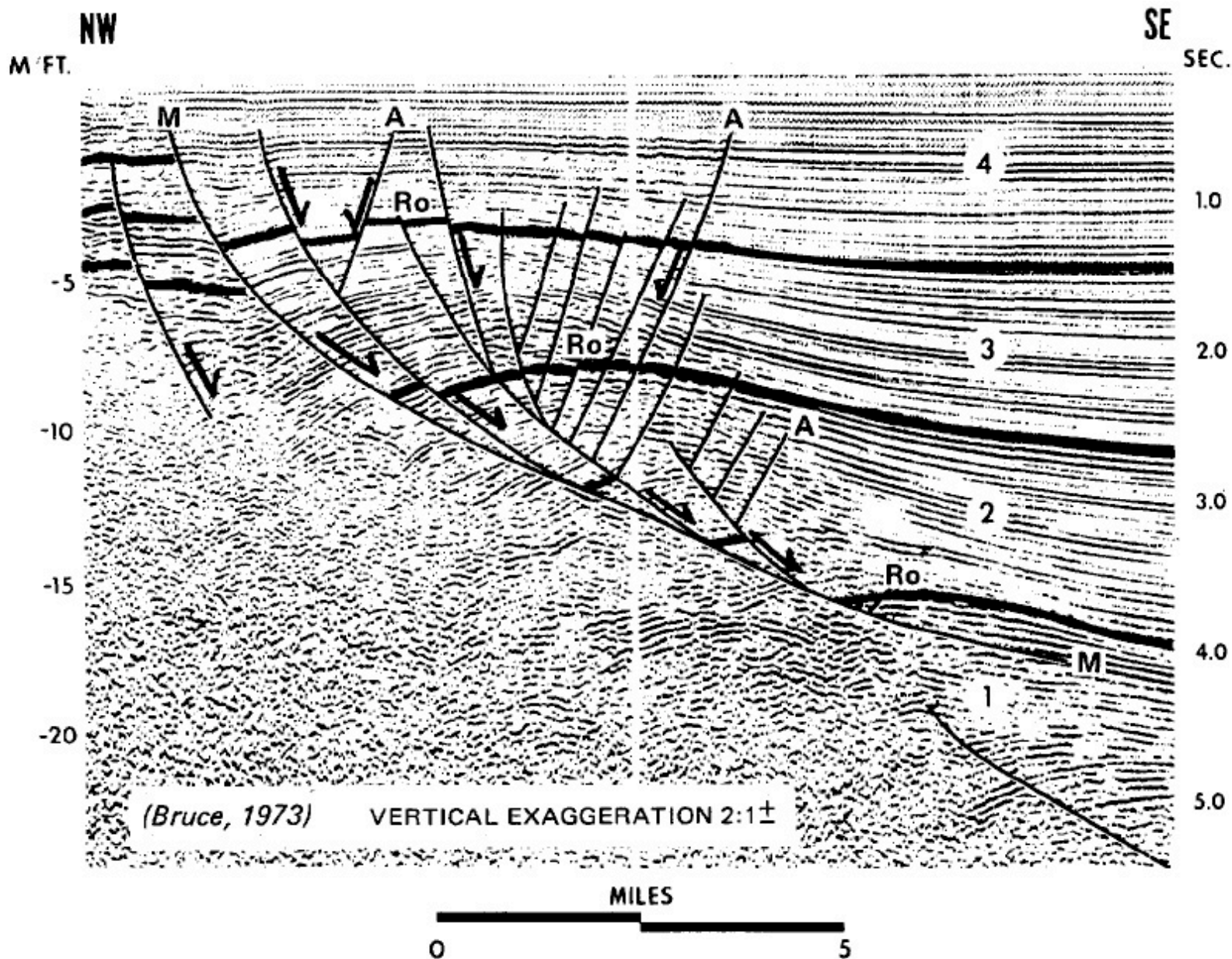


FIG. 21—Syn depositional, detached normal fault assemblage in Tertiary sediments of south Texas part of Gulf Coast basin. *M*, master down-to-basin synthetic fault. *A*, antithetic faults. *Ro*, rollover anticline with basinward migration of crest at depth. Northwest-dipping events beneath *M* fault would migrate to right to form part of rollover anticline. Some assemblages have only synthetic faults.

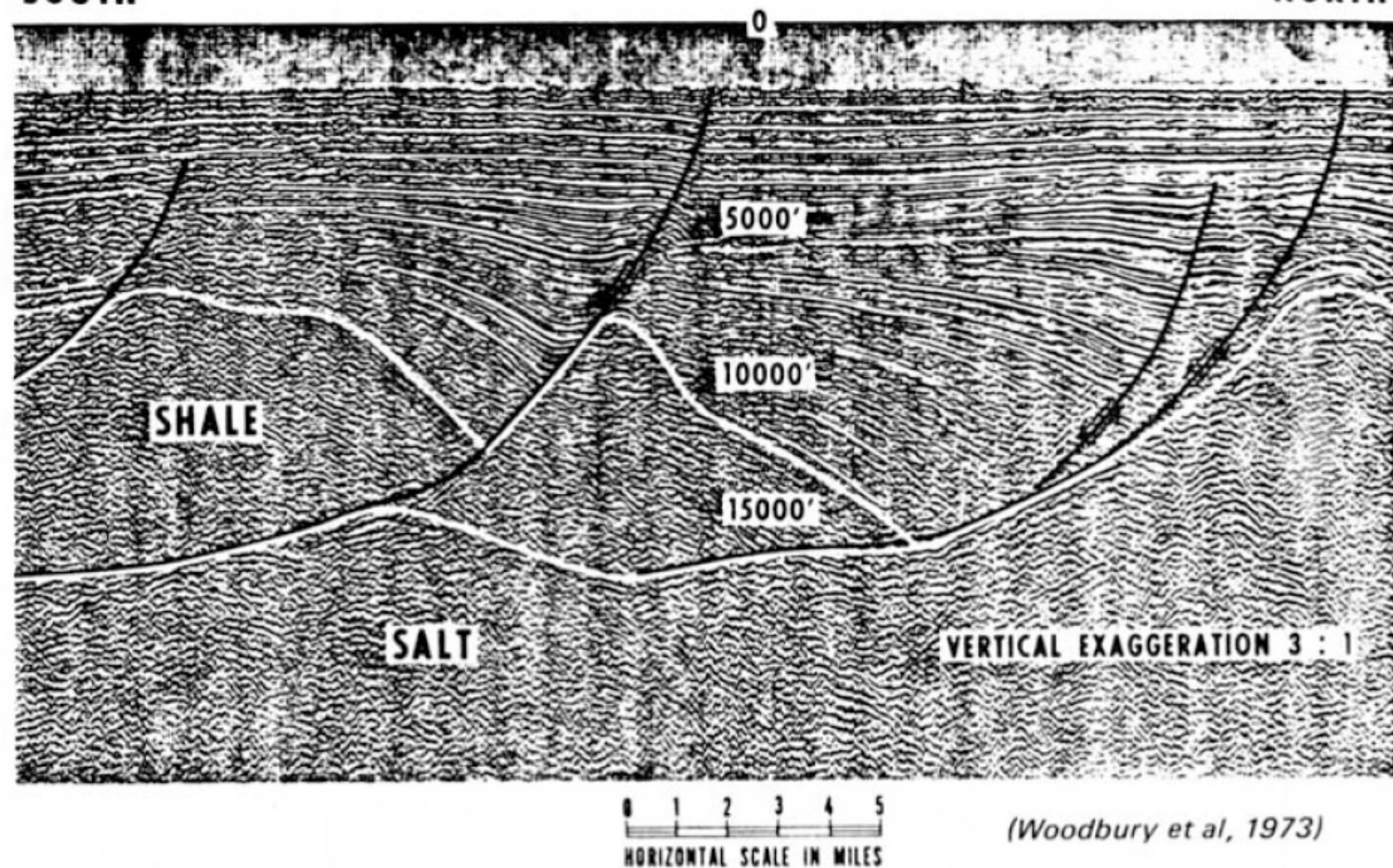
SOUTH**NORTH**

FIG. 22—Syndepositional detached normal faults in outer continental shelf of Gulf Coast basin, Louisiana and Texas. Faults sole out at salt-overburden contact identified by strong reflector and associated diffractions below 15,000 ft (4,500 m) in center of cross section. Presence of diapiric shale core (center) or salt diapir (right) below upthrown side of large growth fault is reported to be very common in this province (Woodbury et al, 1973).

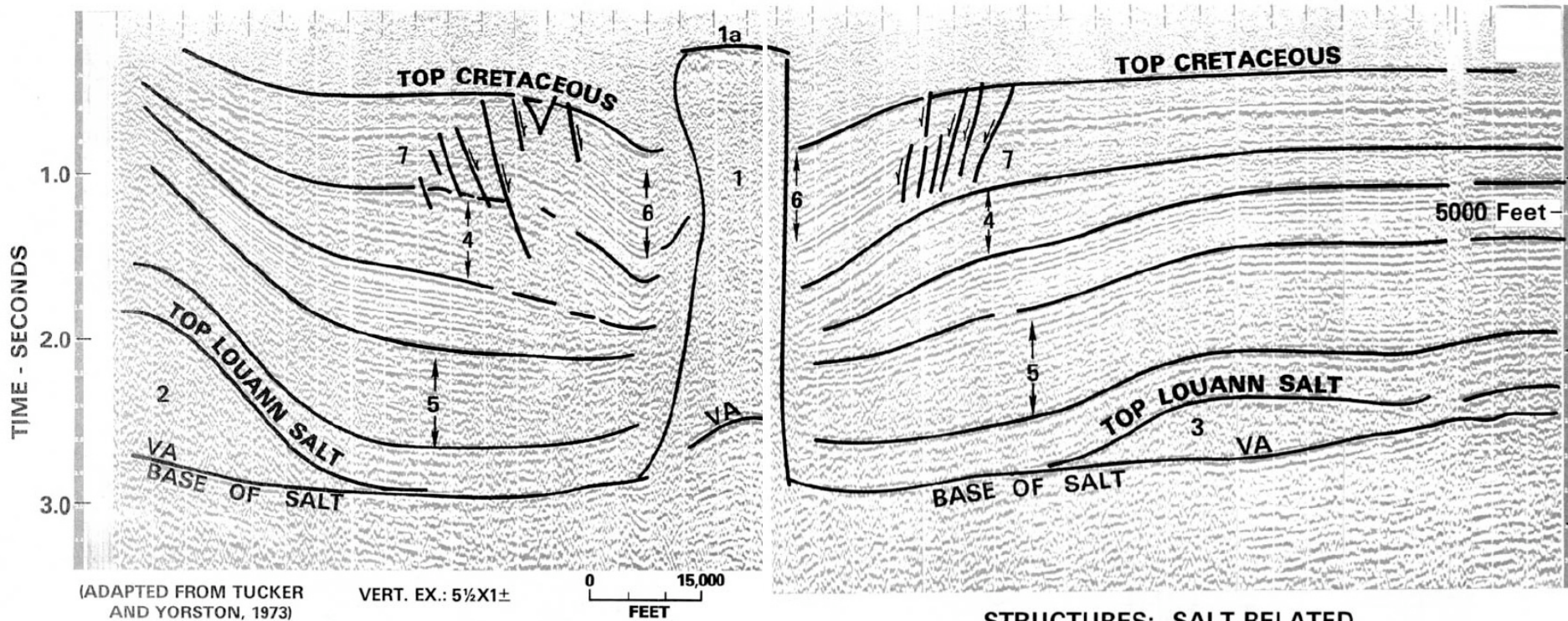


FIG. 26—Seismic section across Hainesville dome (center), east Texas, with structures as indicated. Higher velocity of salt relative to surrounding rocks at velocity anomalies at left and center results in “pull-up” of time section. Velocity anomaly at right shows as depression, because salt here has lower velocity than laterally adjacent rocks.

EXCELLENT

