



Fig. 2. Schematic cross-section of the continental shelf and margin off New York along profile $A \cdot A'$ (Fig. 1). The identification of sedimentary horizons is based mainly on the COST B-2 well and other boreholes in the shelf and margin. The continuity of sedimentary horizons between boreholes is based on the nearby multichannel Line 2 ([4]; Fig. 1). The solid lines in the cross-section are based on prominent seismic horizons summarized in fig. 11 of Schlee et al. [4].

Atlantic Passive Margin



Gulf Coast Passive Margin

Arctic Passive Margin



Other models proposed for the formation of the Amerasian Basin -none except the rotation model are viable for Alaska given the geologic and magneti anomaly constraints

Rifting models for Amerasian Basin summarized by Lawver and Scotese (1990) with specific predictions for rift versus transform origin of marginsand the geologic matches of margins. Base map IBCAO (2000)



Slide courtesv EL Miller. Stanford





Cambrian Limestone









113 Puget Sound Washington





Section B, lower part of succession (from King, 1948).











Williston Basin: Paleocene Sentinel Butte Fm, Roosevelt NP North Dakota





Michigan Basin: Ordovician Queenston Shale, Cheltenham Badlands, Ontario



Patriot Coal Mine, Illinois Basin

http://www.patriotcoal.com/

Albany Shale Map Illinois Basin



Hudson Bay Basin



U Ordovican Beach Sandstones





Typical Hudson Bay Geologist in Field Gear



Two Models for Cratonic Basins



Two Models for Cratonic Basins: How do we test these models?



There is nothing to image with seismic data: no faults. The eclogite keel is too deep to reach by drilling.



Note that the scale of the cratonic basins is realatively small.

Basin Models

- Rift basins: faulting in crust
- Fore-arc basins: slab rollback
- Back-arc basins: slab rollback
- Foredeep basins: response to thrust loads
- Craton basins
 - ?deep crustal faults-not imaged on seismic
 - ?elevated Moho, mantle dymics-not imaged
 - ?cooling lithosphere anomaly-?what evidence



Sloss 1963



Vail et al 1977





Vail et al 1977 sea level curve

Sloss 1963 Sequences

APPALACHIAN

MIOGEOSYNCLINE

QUATERNARY

CRETACEOUS

TERTIARY

JURASSIC

TRIASSIC

PERMIAN

MISSISSIPPIAN

DEVONIAN

SILURIAN

ORDOVICIAN

CAMBRIAN

DOFCAM

TEJAS

PENNSYLVANIAN ABSAROKA

CORDILLERAN

MIOGEOSYNCLINE

ZUNI

KASKASKIA

TIPPECANOE

SAUK





ASSUMPTIONS:

- 1. Stable Continental Platforms
- 2. First and Second Order Cycles represent changes sea floor spreading rate
- 3. Rapid sea floor spreading reduces ocean basin volume, displacing sea water onto shelves

PROBLEMS WITH EUSTACY:

- 1. Continental Platforms Do Not Appear to be Stable in simulations of mantle flow
- 2. First and Second Order Cycles may represent changes in dynamic topography as well as changes in sea floor spreading.
- 3. "Say Goodby to Global Eustacy"-Jerry Mitrovica, AGU Fall Meeting 2008

http://www.agu.org/meetings/fm09/lectures/lecture_videos/T34A.shtml



The Lifespan Of Sedimentary Basin

Woodcock 2005



Figure 2. Typical fates of different classes of sedimentary basin through time, plotted on logarithmic scale. Small Italics show basement to each class. Large Italics show deformational or thermal consequent fates of each class. Estimates of basin heat flow are from compilation by Allen and Allen (1990).



Figure 3. Average life span of basin classes used to calibrate Wilson Cycle of ocean opening and closing. Note linear time scale. Dashed lines are sequential links between basin classes that help to constrain lengths of cycle phases. A continent rifts when it breaks up

6. The continent erodes, thinning the crust

5. As two continents collide orogeny thickens the crust and building mountains



Wilson cycle

2. As spreading continues an ocean opens, passive margin cools and sediments accumulate

3. Convergence begins; an oceanic plate subducts, creating a volcanic chain at an active margin

4. Terraine accretion-from the sedimentary wedge welds material to the continent

