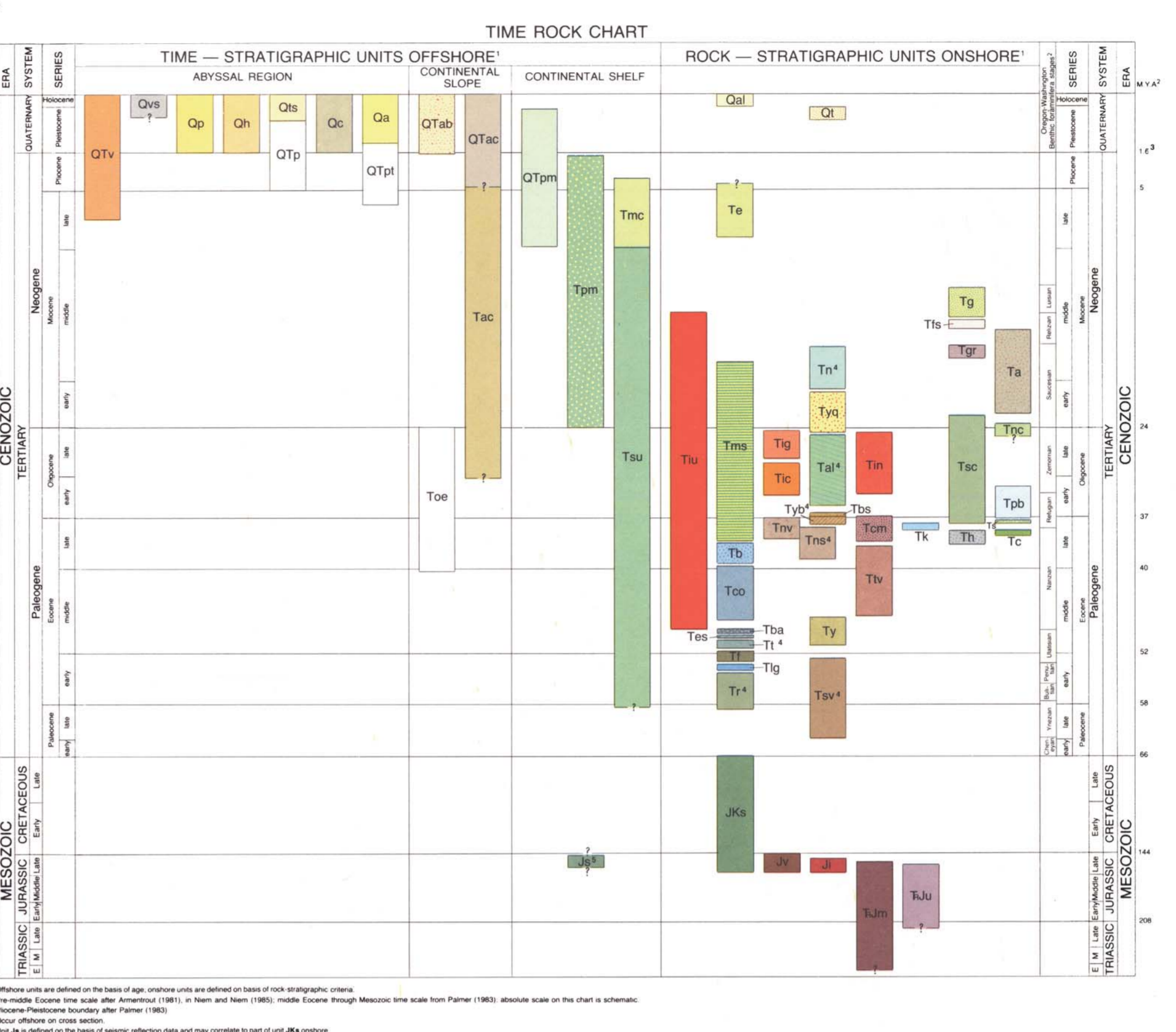
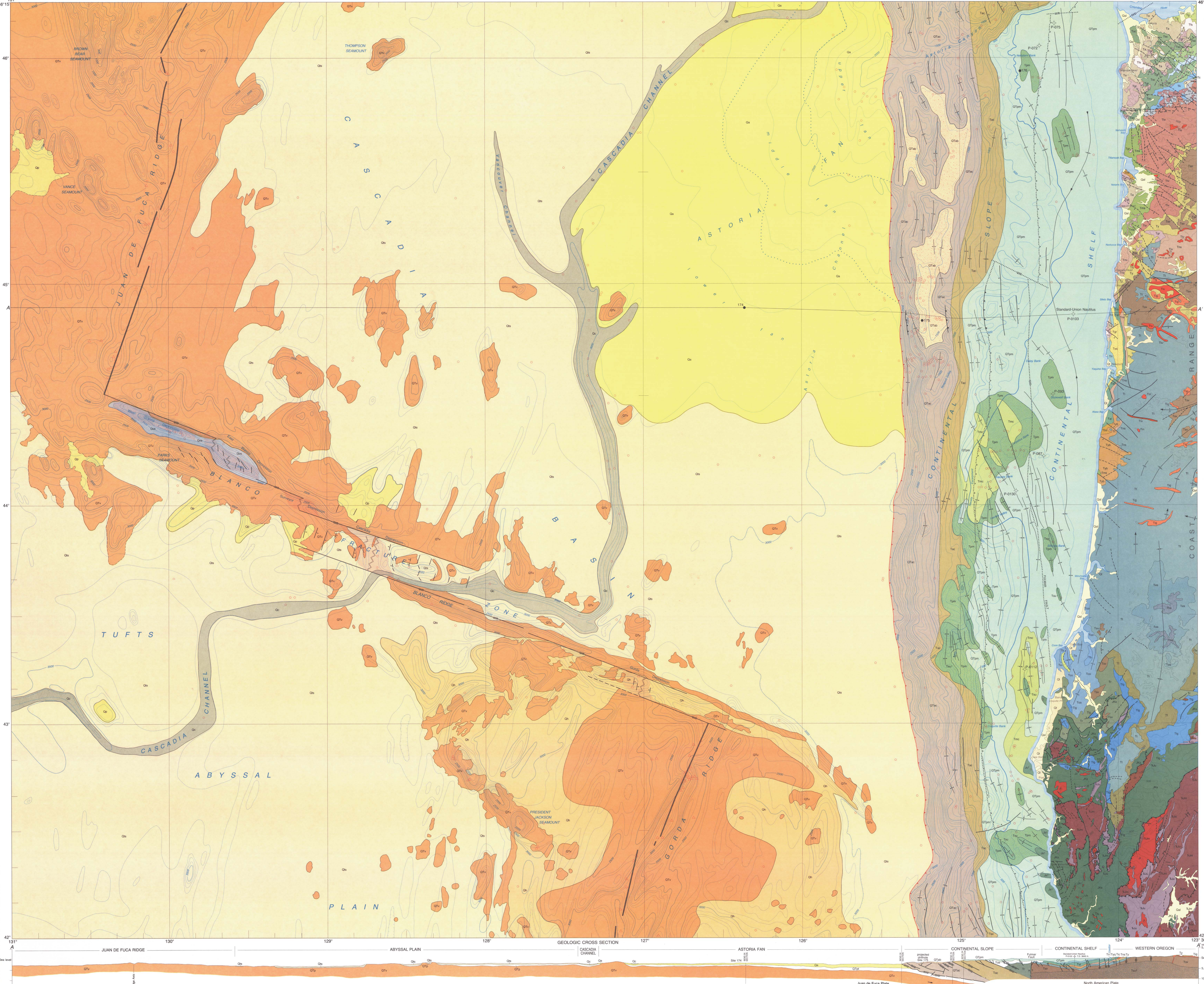


GEOLOGIC MAP OF THE OCEAN FLOOR OFF OREGON AND THE ADJACENT CONTINENTAL MARGIN

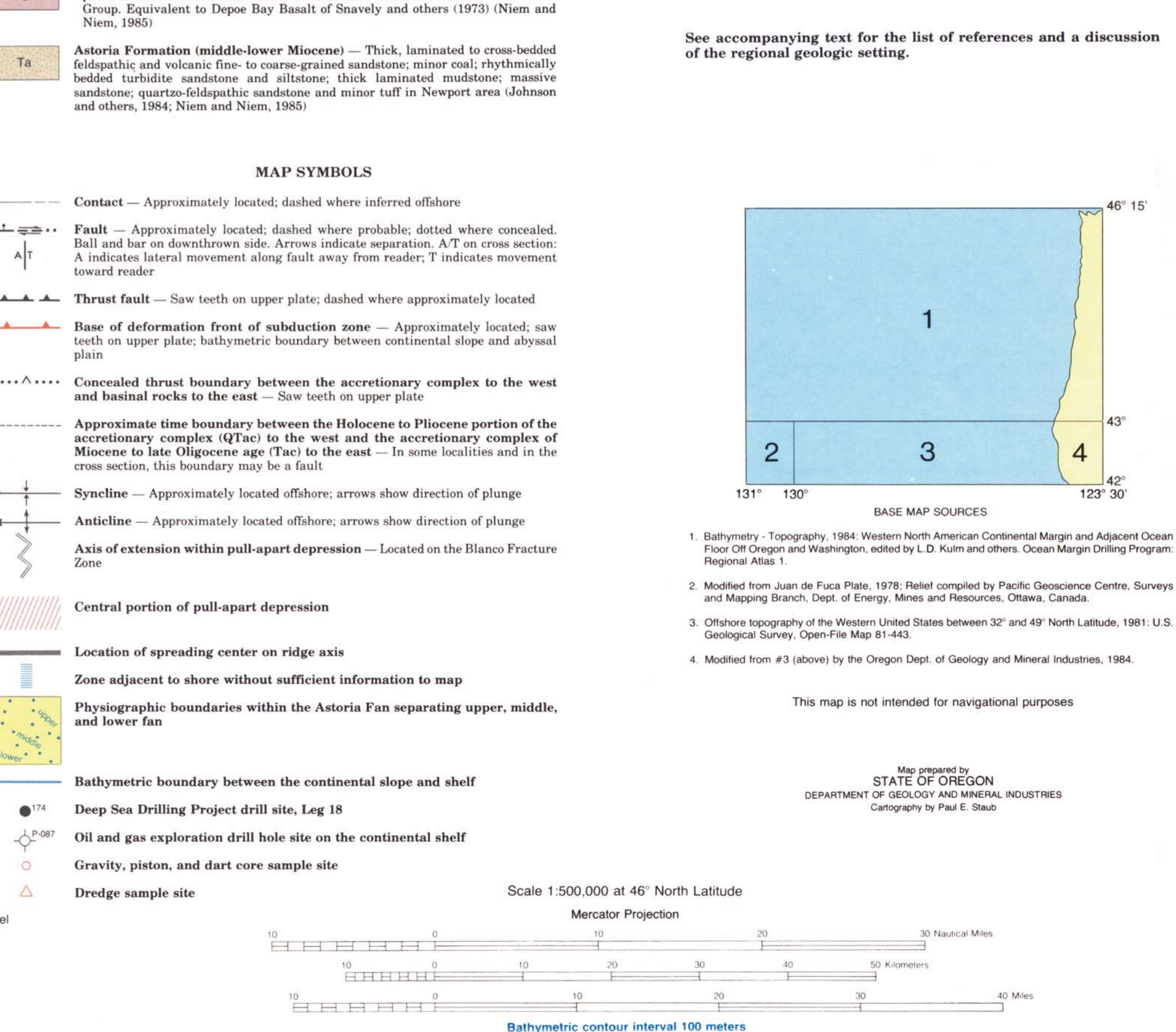
STATE OF OREGON
DEPARTMENT OF GEOLOGY AND MINERAL INDUSTRIES
DONALD A. HULL, STATE GEOLOGIST

By Carolyn P. Peterson and LaVerne D. Kulin, College of Oceanography, Oregon State University, and Jerry J. Gray, Oregon Department of Geology and Mineral Industries

GMS-42
Geologic Map of the Ocean Floor Off Oregon and the Adjacent Continental Margin
By Carolyn P. Peterson, LaVerne D. Kulin, and Jerry J. Gray
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- EXPLANATION**
- OFFSHORE TIME STRATIGRAPHIC UNITS**
 - Juan de Fuca Ridge - Blaine Formation Zone - Gorda Ridge - Abyssal Plain**
 - Manzanita-entirely clay-like (Blaine-Blaine Phosconite?)** - Light and dark brown loamy clay with all forms composed of variably proportioned fine-grained argillaceous, argillaceous, calcareous, siliceous, and siliceous detritus, and some argillaceous detritus. Contains rounded to sub-angular pebbles of sandstone and siltstone. Interbedded with thin, silty, calcareous, and siliceous siltstone. (Nelson and Nunn, 1960)
 - Almond plain turbidite (Blaine-Blaine Phosconite?)** - This unit is the most light-colored and is composed of fine-grained, silty, calcareous, and siliceous detritus. It is a fine-grained turbidite sand with thin, silty, calcareous, and siliceous siltstone. (Nelson and Nunn, 1960)
 - Astoria Fan (Blaine-Blaine Phosconite?)** - Subdivided into three physiographic provinces: upper, middle, and lower. These are thin-bedded, silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Hemipelagic sand (Blaine-Blaine Phosconite?)** - Homogeneous, poorly sorted, silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Pelagic zone and clay (Blaine-Blaine Phosconite?)** - Foraminiferal zone and clay. (Nelson and Nunn, 1960)
 - Cascadia Channel (Blaine-Blaine Phosconite?)** - Rhythmic, interbedded thick and thin beds of silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Almond plain silt turbidite and interbedded clay-like Phosconite-Phosconite** - Interbedded with increasing degree of silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Mid-sea ridge basalt (Blaine-Blaine Phosconite?)** - Thick, pillow lava flow and silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Almond plain silt turbidite (Blaine-Blaine Phosconite?)** - This unit is the most light-colored and is composed of fine-grained, silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Accretionary basin (Blaine-Blaine Phosconite?)** - Interbedded pelagic clay and siltstone. (Nelson and Nunn, 1960)
 - Accretionary complex (Blaine-Blaine Phosconite?)** - Almond plain and fan sediments that have been accreted to the lower continental slope. (Nelson and Nunn, 1960)
 - Almond outcrop and broken formation (Blaine-Blaine Phosconite?)** - May be split, based on lithology and degree of interbedding. (Nelson and Nunn, 1960)
 - Marine sedimentary rocks (Blaine-Blaine Phosconite?)** - Tuffaceous marine siltstone and siltstone. (Nelson and Nunn, 1960)
 - Sedimentary rocks (Blaine-Blaine Phosconite?)** - Shattered assemblage of sedimentary rocks. (Nelson and Nunn, 1960)
 - ONSHORE ROCK STRATIGRAPHIC UNITS**
 - Central Range - Western Coast Range - Klamath Mountains**
 - Almond (Blaine-Blaine Phosconite?)** - Light gray to olive gray, well-sorted, silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Terrace (Blaine-Blaine Phosconite?)** - Elevated marine terrace of unconsolidated silt, sand, and gravel. (Nelson and Nunn, 1960)
 - Esperanza (Blaine-Blaine Phosconite?)** - Facilitative complex of siltstone and sandstone. (Nelson and Nunn, 1960)
 - Great Creek Formation (Blaine-Blaine Phosconite?)** - Marine and terrigenous, thick, cross-bedded, silty, calcareous, and siliceous detritus. (Nelson and Nunn, 1960)
 - Columbia River Basalt Group**
 - Frenchman Springs Member of the Wasnappe Basalt (middle Miocene)** - Basaltic and andesitic flows and tuffs. (Nelson and Nunn, 1960)
 - Grande Ronde Basalt (middle Miocene)** - Basaltic and andesitic flows and tuffs. (Nelson and Nunn, 1960)
 - Astoria Formation (middle-lower Miocene)** - Thick, laminated, cross-bedded turbidite sandstone and siltstone. (Nelson and Nunn, 1960)



The Geology of the Ocean Floor Off Oregon and the Adjacent Continental Margin

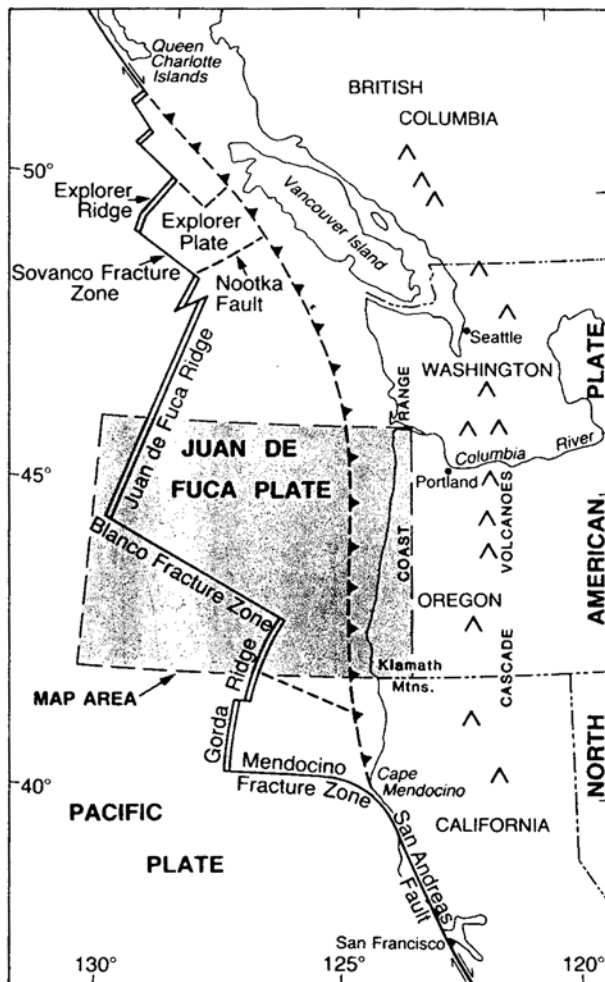
By Carolyn P. Peterson and LaVerne D. Kulm, Oregon State University, and
Jerry J. Gray, Oregon Department of Geology and Mineral Industries

Introduction

The geologic map was constructed from a large number of published sources, including the papers, maps, charts, and side-scan sonar mosaics cited in the *Geologic Bibliography and Index Maps of the Ocean Floor Off Oregon and the Adjacent Continental Margin* (Peterson and others, 1985), which is a companion publication to this one; the atlas of the *Western North American Continental Margin and Adjacent Ocean Floor Off Oregon and Washington* (Kulm and others, 1984b); and *Atlas of the Exclusive Economic Zone, Western Conterminous United States* (EEZ Scan 84 Scientific Staff, 1986). The major sources of published and unpublished data used in this map are identified in the sections entitled "Acknowledgments" and "References." A substantial amount of data resides in the College of Oceanography archives at Oregon State University (OSU) and in the Office of Pacific Marine Geology, U.S. Geological Survey (USGS), Menlo Park, California. The base map, which includes the offshore bathymetry and onshore topography, was modified from the map published by Connard and others (1984) for western Oregon and the adjacent seafloor. Although more accurate bathymetric maps have been made through SeaBeam swath mapping of portions of the continental slope and the spreading Gorda and Juan de Fuca Ridges, they have not yet been published for the Oregon region. Nevertheless, selected structural information and trends of tectonic features derived from unpublished SeaBeam data were incorporated into this geologic map. These include the major faults and pull-apart basins (e.g., Cascadia Depression) of the Blanco Fracture Zone and the trends of axial zones of the spreading ridges (Embley and others, 1986). Discrepancies between the two bathymetric data sets have been resolved to achieve a best fit of these data.

Offshore Geology

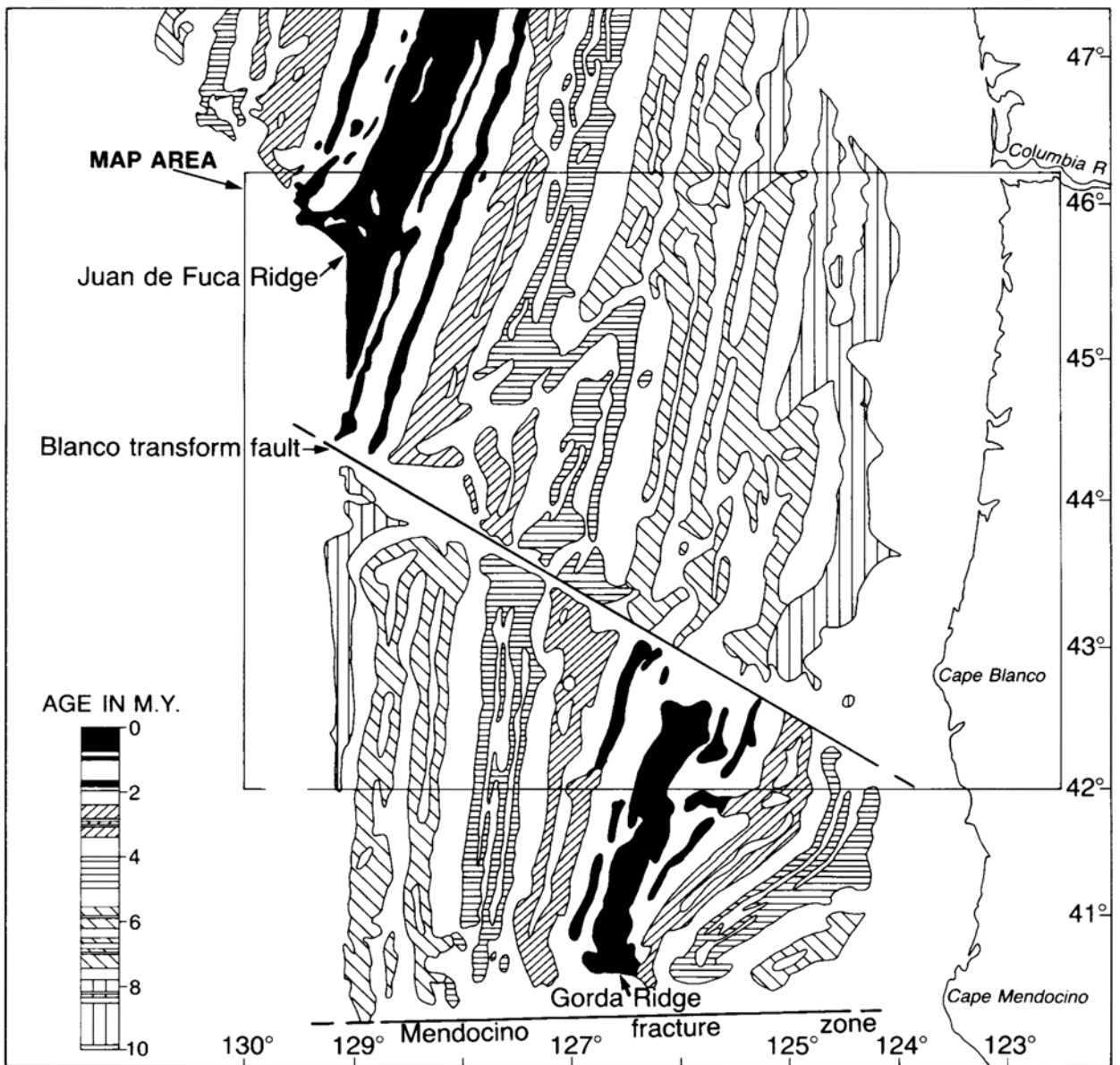
Gravity, piston, and dart core samples (sample locations shown on the map as open circles), dredge samples (open triangles), data from Deep Sea Drilling Project (DSDP) Site 174, seismic reflection records, and observations made from the submersible *Alvin* were used to define the abyssal plain stratigraphic units and boundaries. The geologic units of the abyssal regions are (1) basalt (QTv) comprising the young spreading ridges and seamounts and underlying the Cascadia Basin; (2) manganese-enriched volcanic silts of the East and West Blanco Depressions (Qvs); (3) abyssal plain sand-silt turbidites and hemipelagic mudstones (Qts); (4) sand turbidites of the Astoria Fan (Qa), which are prograding seaward over the abyssal plain deposits (QTpt on the cross section); and (5) relatively thin hemipelagic (Qh) and pelagic (Qp) deposits draped over the volcanic topography (seamounts) of the Cascadia Basin and portions of the mountainous spreading ridges. These volcanic terrains are covered by either hemipelagic deposits (i.e., Qh on the Gorda Ridge) or pelagic deposits (i.e., Qp on the Juan de Fuca Ridge) shortly after their formation, depending upon their proximity to the continent. Hemipelagic deposits consist largely of terrigenous silts and clays. With increasing distance from the continent, the hemipelagic deposits grade into pelagic sediments comprised mainly of foraminiferal and radiolarian ooze with minor clay (Duncan, 1968; Phipps, 1974). When the warm buoyant oceanic crust cools and subsides on either side of the spreading centers, these hemipelagic/pelagic deposits are covered by terrigenous turbidites. Core lithologies, depositional patterns, and concepts of turbidity current deposition for the Cascadia Basin/Tufts Abyssal Plain, Cascadia Channel, and Blanco Fracture Zone, as described by Duncan (1968), Griggs and Kulm (1970), and Ibach (1981), respectively, were used to define and extrapolate the boundaries of the silt and sand turbidites of late Pleistocene age (Qts) on the geologic map. Terrigenous turbidites may directly overlie volcanic rocks where the tectonic setting allows turbidity currents access to the newly formed oceanic crust of the pull-apart basins (i.e., unit Qts in the Cascadia Depression; Embley and others, 1986). The sandy turbidites (Qc) of the Cascadia Channel indicate that the channel has served as a conduit for turbidity currents reaching the Blanco Fracture Zone and Tufts Abyssal Plain (Griggs and Kulm, 1970, 1973; Duncan and



Location map of the Juan de Fuca plate and the Oregon continental margin (from Riddiough, 1984).

Kulm, 1970). Portions of the volcanic terrain of the Blanco Ridge were uplifted during late Pleistocene time, isolating terrigenous turbidites on these topographic highs (Duncan, 1968; Duncan and Kulm, 1970; Ibach, 1981).

A variety of geological and geophysical data were used to define the stratigraphic units and the structures of the continental shelf and slope. Data from six oil and gas exploration drill holes (Snavey and others, 1982; Kulm and others, 1984a; Peterson and others, 1984), DSDP Sites 175 and 176 (Kulm and others, 1973b,c), and approximately 250 dredge and core samples, concentrated largely in the vicinity of submarine banks (Maloney, 1965; Carlson, 1968; Fowler and others, 1971; Kulm and Fowler, 1974; Kulm and others, 1984a) were used to define the stratigraphic units on the continental shelf. Single-channel seismic-reflection records (e.g., Mackay, 1969; Muehlberg, 1971; Spigai, 1971; Kulm and Fowler, 1974; Clarke and others, 1985) and multichannel records (Snavey and others, 1977, 1980, 1985, 1986; Mann and Snavey, 1984) were used to extrapolate the geologic units to areas of the shelf where no samples were available and to delineate the major structures. A major right-lateral strike-slip fault (Fulmar Fault) is inferred from oil and gas exploration drill holes and seismic and magnetic data on the continental shelf. The major movement on the Fulmar Fault occurred prior to the deposition of strata of late Eocene to Oligocene age (Toe) that overlap



Magnetic anomaly map (modified from Silver, 1971).

the fault (Snively and others, 1977, 1980, 1982; Snively, 1986). The fold pattern in the sediments of the continental shelf (**QTpm**) and slope (**QTac**, **Tac**) trends in a north-south direction, whereas structures in the older strata (**Tpm**) have a more northwesterly trend. Interconnecting shallow synclines occurring between the shore and the submarine banks near the outer edge of the shelf contain largely siltstones and sandstones of Pleistocene to late Miocene age (**QTpm**). Some of the most striking folded and faulted areas occur beneath the submarine banks (e.g., Nehalem, Heceta, and Coquille Banks), which have been uplifted as much as 1 kilometer (km) since late Miocene time (Kulm and Fowler, 1974). The most structurally complex area of the inner shelf lies between Coos Bay and the Rogue River. Pre-Quaternary sediments (**Tmc**) tentatively correlated with onshore sediments (**Te**; Fowler and others, 1971) crop out on the seafloor or are covered by a thin veneer of Quaternary sediment. Pre-Tertiary strata (**Js**) form an irregular acoustic basement that shallows near the shore in the vicinity of Cape Blanco and in isolated outcrops southward to the California border.

Lithologic and tectonic data derived from DSDP Site 175 (Kulm and others, 1973b), the structural and stratigraphic data from a 1984 study using the submersible *Alvin* (Kulm and others, 1986), deep-towed camera surveys, and dated punch cores and dredges were used to identify the Holocene to Pliocene portion of the ac-

cretory complex (i.e., unit **QTac**, accreted strata comprised largely of a seaward-verging thrust sequence of mudstones and sandstones), which is exposed in canyons and slump scars on the lower and middle continental slope. On the geologic map, the accretory complex has been divided into a younger unit of Holocene to Pliocene age (**QTac**) and an older unit of Miocene to late Oligocene age (**Tac**). Older rocks of the accretory complex of Miocene to late Oligocene age (unit **Tac** on cross section) are most likely situated landward of unit **QTac** on the upper continental slope (Snively and others, 1980). Overlying Holocene to Pleistocene basinal deposits (**QTab**) are recognized in Site 175 and in seismic-reflection records made over the continental slope. In the areas of the continental slope that were not surveyed or sampled, multichannel seismic-reflection records were used to define the accretory complex; these relatively thin basin deposits are clearly outlined in seismic-reflection records. The older and more landward subsurface portions of the inferred accretory complex (**Tac**) are unsampled. The boundary between the accreted strata (**Tac**) of Miocene to late Oligocene age and the Eocene rocks occurring within units **Toe**, **Tsv**, and **Tr** on the cross section, as shown on the geologic map, was inferred from the noticeable seaward disruption of strata shown in multichannel seismic records made across the outermost continental shelf.

Observations from the submersible *Alvin* and *Alvin*-towed seis-

mic-reflection records made off central Oregon show that the most recent thrust boundary between the subducting Juan de Fuca Plate and the North American Plate (lowermost continental slope) is a deformation front characterized by a series of low-lying benches consisting generally of seaward-verging strata (Kulm and others, 1986). While landward-verging strata are observed in some submersible-towed and surface-towed multichannel seismic records, they appear to be confined to the shallower portions of the overall underthrusting sedimentary sequences.

The generalized geologic cross section (A-A') was derived and simplified from several data sources: abyssal plain from Griggs and Kulm (1973), Kulm and others (1973a), and OSU data files; continental slope from Kulm and others (1973a), von Huene and Kulm (1973), Snively and others (1980, 1986), and Kulm and others (1986); continental shelf from Snively and others (1980); and western Oregon from Snively and others (1976b, 1980) and Johnson and others (1984).

Onshore Geology

The geologic map of western Oregon is simplified from several published sources. Compilation of published and unpublished maps by Wells and others (1983), Johnson and others (1984), and Niem and Niem (1985) are the primary sources of the map north of 43° N. latitude to the Columbia River. County geologic maps (Ramp, 1972; Baldwin and others, 1973; Ramp and others, 1977; Ramp and Peterson, 1979) were used for the compilation south of 43° N. latitude to the Oregon-California border. Tertiary map units are shown on the geologic map as they are given in the source maps. The Mesozoic rock units have been simplified and combined into five units: sedimentary, volcanic, igneous intrusive, ultramafic, and metamorphic rocks.

Acknowledgments

Several individuals kindly contributed diagrams and data of both unpublished and published information for the compilation of the geologic map. The locations of the axes of the spreading Gorda and Juan de Fuca Ridges and the structure of the pull-apart basins of the Blanco Fracture Zone were provided by Robert Embley and Steven Hammond, National Oceanic and Atmospheric Administration (NOAA), from the SeaBeam bathymetry of these features. Sam Clarke and Parke D. Snively, Jr., USGS, contributed information and valuable discussion on the structure and stratigraphy of the continental shelf and slope off Oregon. Discussions with Alan R. Niem and Wendy A. Niem, OSU; Parke D. Snively, Jr., USGS; and John D. Beaulieu and Beverly F. Vogt, Oregon Department of Geology and Mineral Industries (DOGAMI), about the onshore structure and stratigraphy were very helpful in compiling the map. The descriptions and locations of recently acquired dredge samples from the Gorda Ridge were kindly supplied by Martin Fisk, OSU. Elizabeth Asbury, OSU, produced the computer output of the cores and dredges displayed on the map. Margaret Stribling, OSU, assisted in compiling both onshore and offshore geology. The authors gratefully acknowledge review by Wendy A. Niem and Alan R. Niem, OSU; Parke D. Snively, Jr., and Sam Clarke, USGS; Robert Embley, NOAA; and Donald A. Hull and Len Ramp, DOGAMI. The authors bear the ultimate responsibility for this map. This study was funded by the U.S. Department of the Interior—Minerals Management Service through Cooperative Agreement 14-12-0001-30223.

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Geologic Map of the Ocean Floor Off Oregon and the Adjacent Continental Margin